

Earthquake and Fault Studies by Moment Tensor Inversion and Events Relocation

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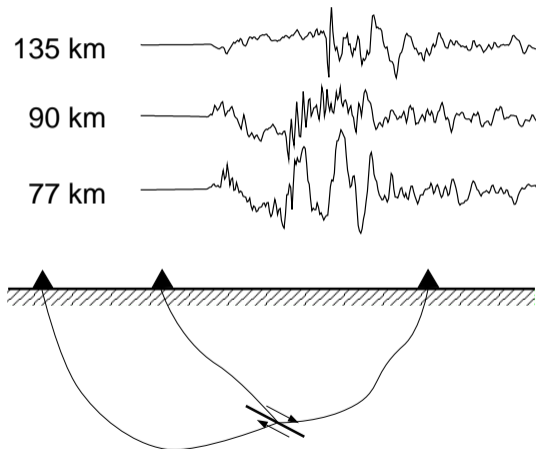
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¹Zhu and Helmberger (1996); Zhu and Ben-Zion (2013); Zhu and Zhou (2016); Huang et al. (2018)

Outline

1. Seismic moment tensor inversion for a general seismic source.
2. The Cut-and-Paste (CAP) moment inversion method and gCAP3D.
3. Application of gCAP3D to the 2013 *M*_s 7.0 Lushan earthquake.
4. Seismic and geological evidence of induced earthquakes in TGR.
5. Conclusions.

Moment tensor and the Green's functions



For a DC point seismic source with an impulsive STF,

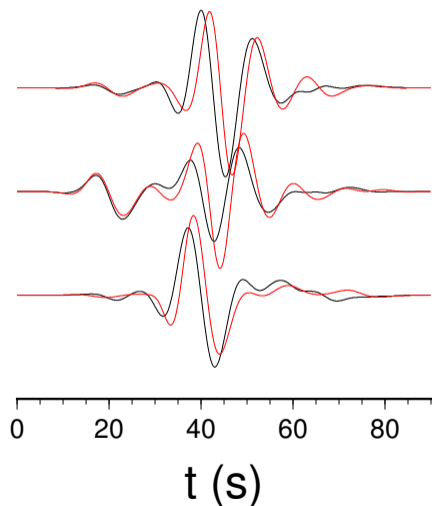
$$u(t) = M_0 \sum_{i=1}^3 A_i(\phi, \delta, \lambda) G_i(t), \quad (1)$$

where $G(t)$'s are Green's functions. For a general point source,

$$u(t) = \sum_{i=1}^3 \sum_{j=1}^3 M_{ij} G_{ij}(t), \quad (2)$$

where \mathbf{M} is the source moment tensor.

Moment tensor inversion

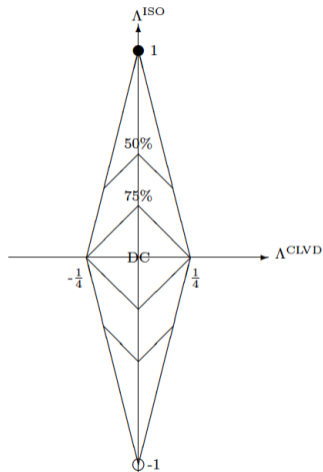


Eq. (2) is a linear function of M_{ij} when the Green's functions are known. But imperfect velocity model and source location/origin time introduce a unknown time shift Δt (Zhao and Helmberger, 1994),

$$u(t) = \sum_{i=1}^3 \sum_{j=1}^3 M_{ij} G_{ij}(t - \Delta t), \quad (3)$$

which makes the moment tensor inversion non-linear.

Parameterization of moment tensor



$$M_{ij} = M_0 D_{ij}, \quad (|\mathbf{D}| = \sqrt{2}) \quad (4)$$

$$D_{ij} = \zeta D_{ij}^{ISO} + \sqrt{1 - \zeta^2} \left(\sqrt{1 - \chi^2} D_{ij}^{DC} + \chi D_{ij}^{CLVD} \right), \quad (5)$$

where

$$D_{ij}^{ISO} = \sqrt{\frac{2}{3}} \delta_{ij}, \quad (6)$$

$$D_{ij}^{DC} = n_i v_j + v_i n_j, \quad (7)$$

$$D_{ij}^{CLVD} = \frac{1}{\sqrt{3}} (2b_i b_j - v_i v_j - n_i n_j), \quad (8)$$

ζ and χ are non-dimensional parameters quantifying the strength of isotropic and CLVD components, respectively (Zhu and Ben-Zion, 2013).

Generalized Cut-and-Paste (gCAP3D) method

gCAP3D (Zhu and Zhou, 2015) uses a grid search to solve Eq. (3) for

$$\mathbf{m} = (\zeta, \chi, \phi, \delta, \lambda)^T. \quad (9)$$

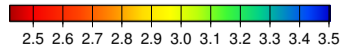
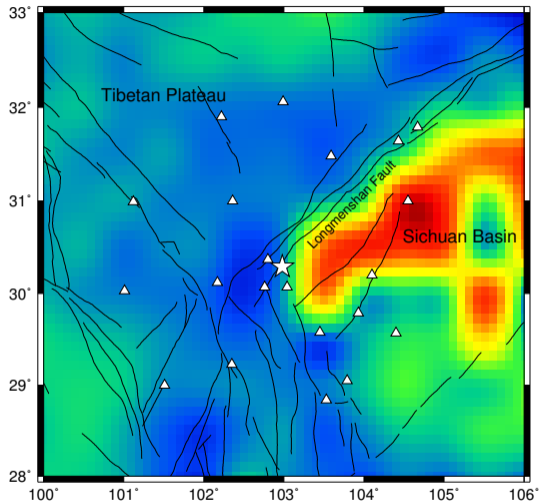
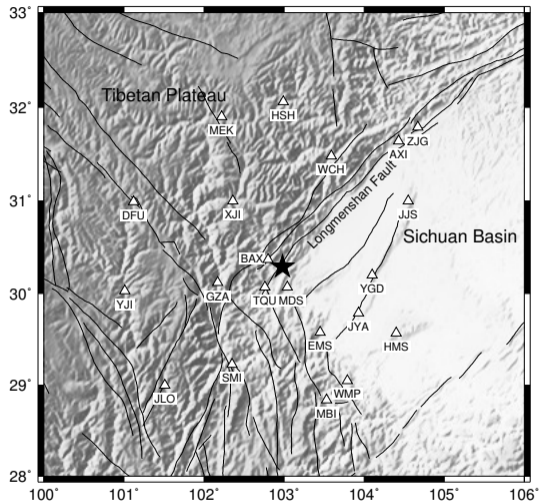
For each possible set of source parameters, it first finds Δt by cross-correlating $u(t)$ and $s(t) = D_{ij}G_{ij}(t)$ and estimates the scalar moment,

$$M_0 = \frac{\|u\|}{\|s\|}. \quad (10)$$

It then calculates the waveform misfit e using the L_2 norm of the difference between observed and predicted waveform,

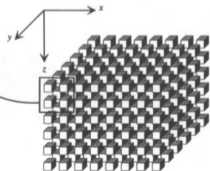
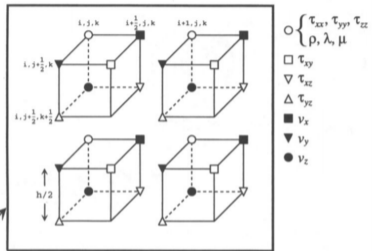
$$E = \sum_{i=1}^{N_s} \left(w^2 \left(\frac{r_i}{r_0} \right)^2 \left(e_i^{PnlZ} + e_i^{PnlR} \right) + \frac{r_i}{r_0} \left(e_i^{RaylZ} + e_i^{RaylR} + e_i^{Love} \right) \right), \quad (11)$$

$$e = \|u(t) - M_0 s(t - \Delta t)\|^2. \quad (12)$$



V_s (km/s) at 5 km

Unit Cell for Staggered-Grid Formulation



- EMOD3D code by R. Graves (1996).
- Staggered grid, 4th-order FD.
- $450 \times 450 \times 150$ km.
- Grid spacing 1 km, $f_{max} = 0.4$ Hz.
- Use the reciprocity principle to reduce the number of FD runs.
- Takes ~ 4 Hrs. per station.



Event 20130420000246 Model and Depth 1D_15

FM 30 54 82 Mw 6.37 E 8.109e+01 589 ERR 2 4 6 ISO 0.00 0.00 CLVD 0.00 0.00

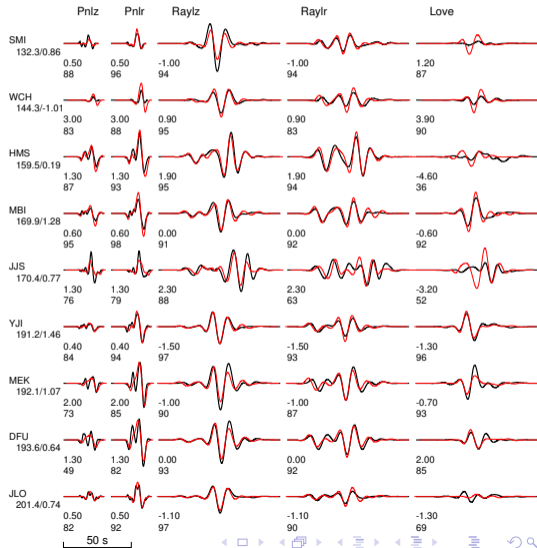
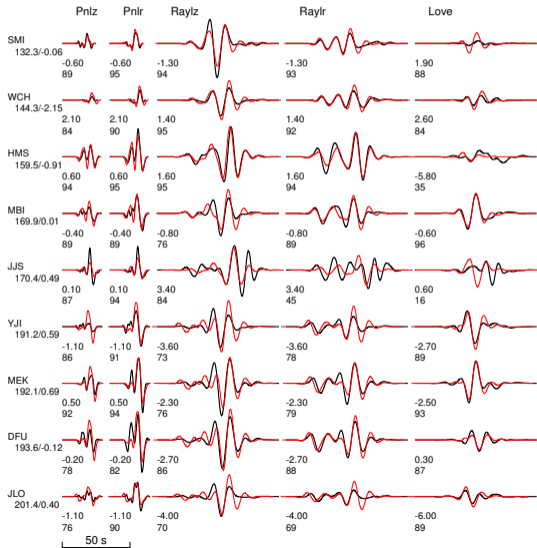
Variance reduction 64.0



Event 20130420000246 Model and Depth 3D_fix

FM 224 41 90 Mw 6.53 E 7.245e+01 590 ERR 2 4 6 ISO 0.00 0.00 CLVD 0.00 0.00

Variance reduction 68.3

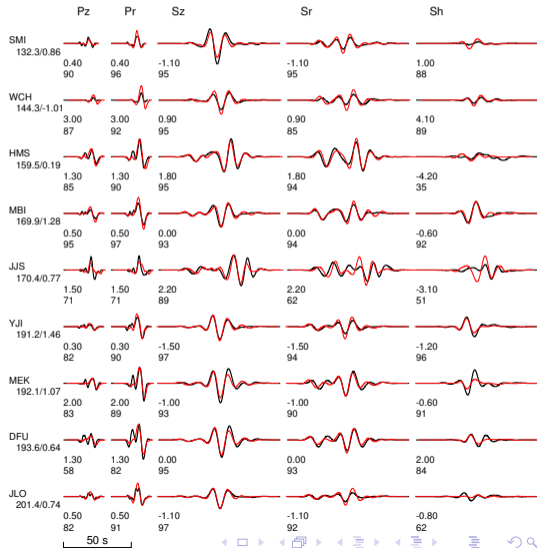
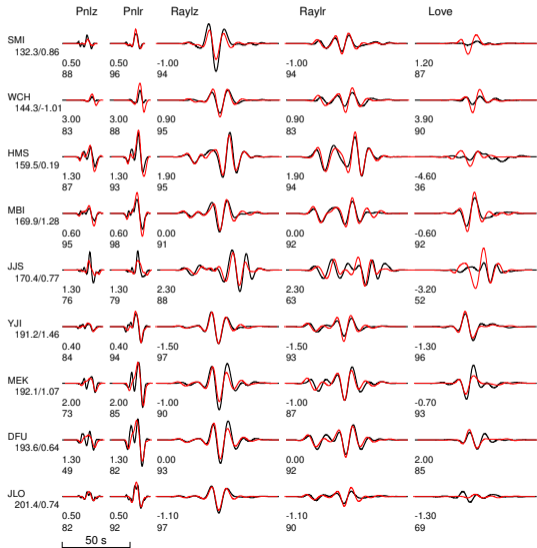


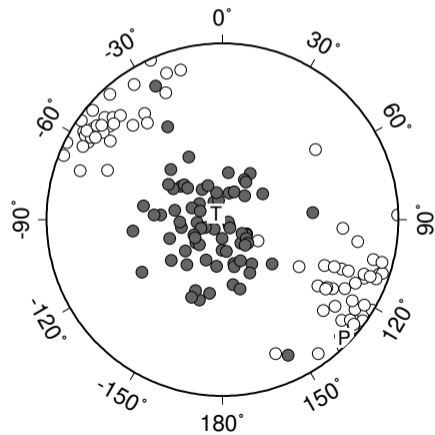
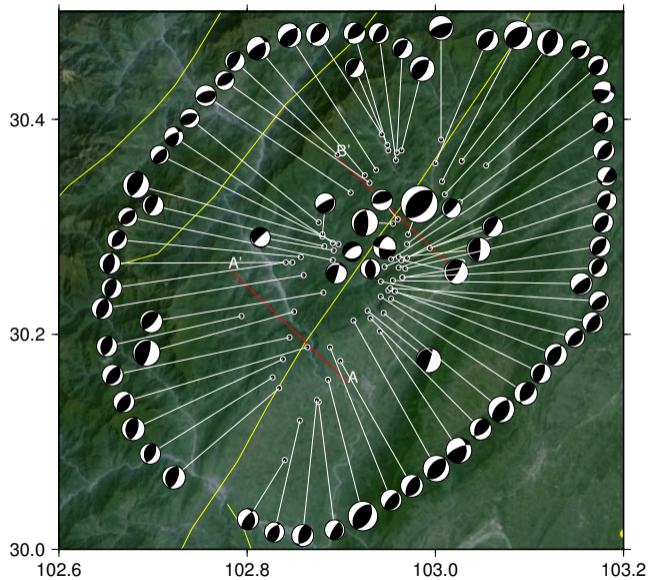


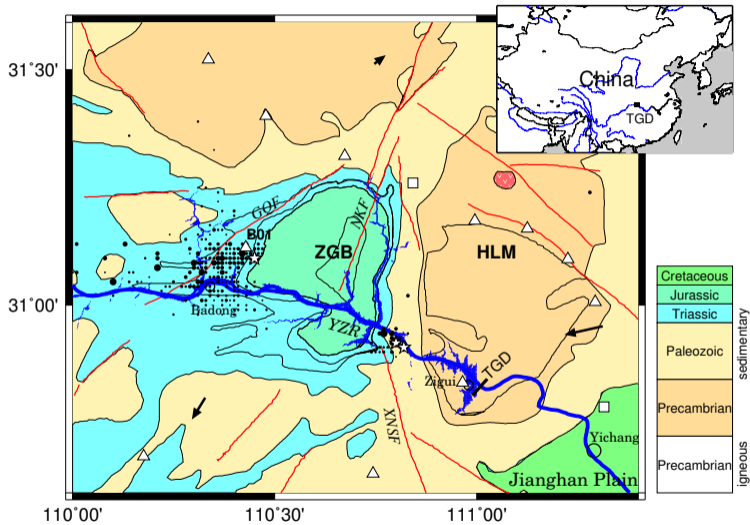
Event 20130420000246 Model and Depth 3D_fix
 FM 224 41 90 Mw 6.53 E 7.245e+01 590 ERR 2 4 6 ISO 0.00 0.00 CLVD 0.00 0.00
 Variance reduction 68.3

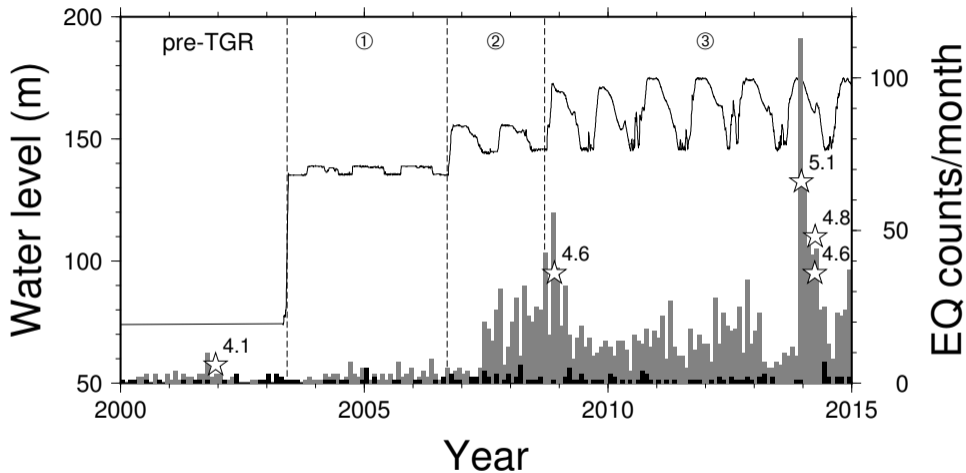


Event 2013mainfmt Model and Depth gecan_fix
 FM 222 42 90 Mw 6.52 E 6.528e+01 590 ERR 2 1 2 ISO 0.28 0.05 CLVD -0.13 0.03
 Variance reduction 71.4



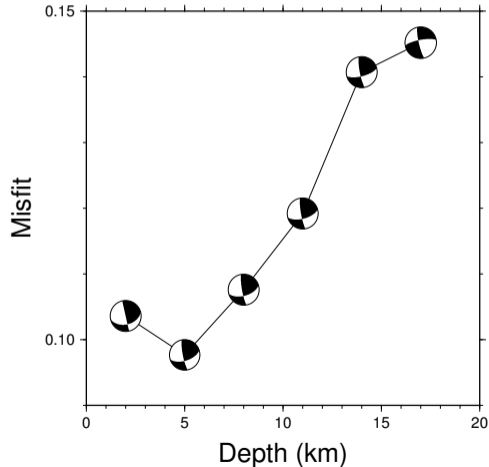
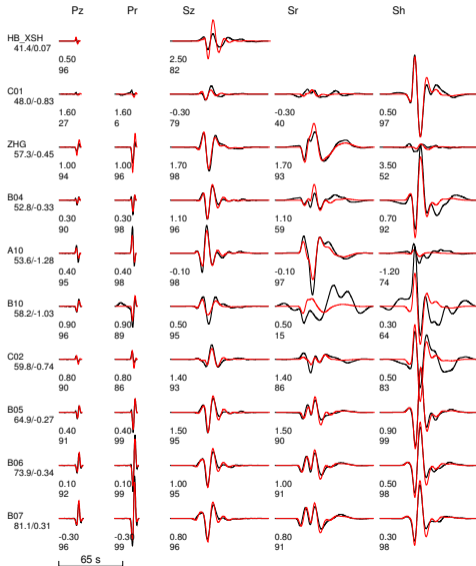


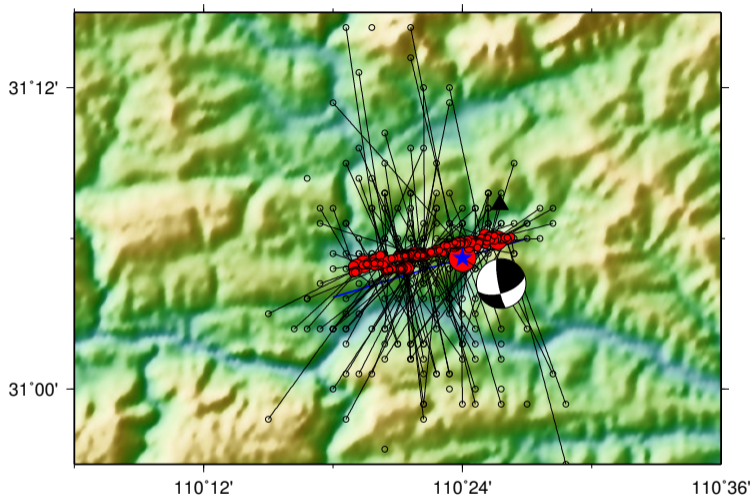




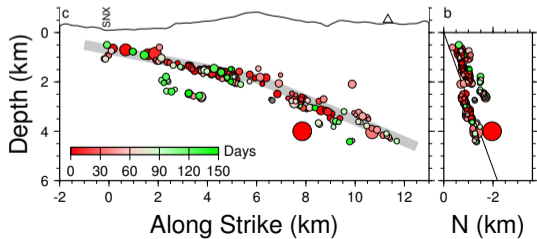
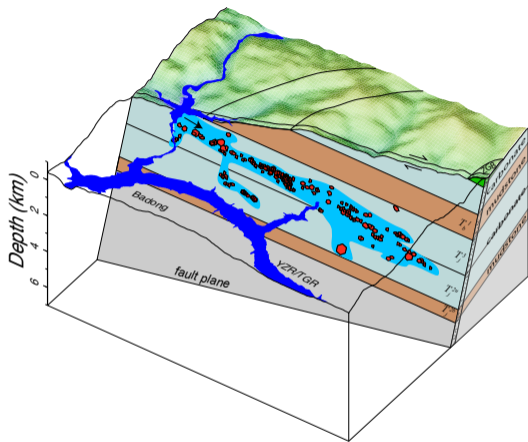
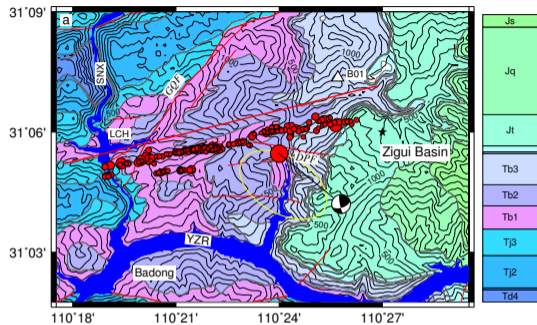


Event 20131216050452 Model and Depth exi_05
 FM 169 80 32 Mw 4.86 E 9.768e-02 3039 ERR 0 1 1 ISO 0.00 0.00 CLVD 0.00 0.00
 Variance reduction 78.3





Out of 219 events with 20 km of the mainshock, 205 were relocated along a line in $N73^\circ$.



Conclusions

1. We developed a method for determining moment tensors using 3D velocity models.
2. The method has been fully automated for real-time moment tensor inversion.
3. We applied the method to the 2013 M_s 7.0 Lushan earthquake sequence and obtained 75 moment tensor solutions ranging from M_w 6.5 to 3.4. The mainshock is a reverse faulting on a plane dipping $40\text{-}47^\circ$ to NW.
4. We also applied the method and HypoDD to the 2013 M 5.1 Badong earthquake sequence in the Three Gorges Reservoir. We found an unmapped fault of ~ 15 km long that is seismically active and has potential to produce a large earthquake of $M > 6.5$. The results suggest that the sequence was induced by water infiltration in a fractured carbonate rock formation from the reservoir.